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# Statistical correlation of the time rate of change of relative vorticity with convective weather phenomena

Writt, Elinor Jane

Monterey, California: U.S. Naval Postgraduate School

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STATISTICAL CORRELATION  
OF THE  
TIME RATE OF CHANGE OF  
RELATIVE VORTICITY WITH  
CONVECTIVE WEATHER PHENOMENA

BY  
ELINOR JANE WRITT

THESIS  
W92

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A STATISTICAL CORRELATION OF THE TIME RATE OF CHANGE OF  
RELATIVE VORTICITY WITH CONVECTIVE WEATHER PHENOMENA

by  
Elinor Jane Writt  
"



Thesis  
W92

THE UNIVERSITY OF CHICAGO  
DEPARTMENT OF THE HISTORY OF ARTS  
AND ARCHITECTURE

OF  
THE HISTORY OF ARTS

A STATISTICAL CORRELATION OF THE TIME RATE OF CHANGE OF  
RELATIVE VORTICITY WITH CONVECTIVE WEATHER PHENOMENA

by  
Elinor Jane Writt  
Lieutenant, United States Navy

Submitted in partial fulfillment  
of the requirements  
for the degree of  
MASTER OF SCIENCE  
IN AEROLOGY

United States Naval Postgraduate School  
Monterey, California  
1951



This work is accepted as fulfilling  
the thesis requirements for the degree of

MASTER OF SCIENCE  
IN AEROLOGY



## PREFACE

This paper represents a statistical evaluation of the relationship between relative vorticity and convective weather phenomena. Specifically, it contains a correlation of both vorticity neglecting wind shear and individual rate of change of vorticity, with convective activity. Data for these determinations was obtained from surface maps covering a two and one-half year period, from January, 1949 to May, 1951.

This work was conducted at the United States Naval Postgraduate School, Monterey, California in the spring of 1951 in partial fulfillment of the requirements for the degree of Master of Science in Aerology.

The author wishes to express gratitude for the assistance of Associate Professor G. J. Haltiner, Department of Aerology, U. S. Naval Postgraduate School, Monterey, California.





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# TABLE OF SYMBOLS AND ABBREVIATIONS

$C$	Relative circulation
$\Delta A$	Element of area
$\delta A$	Element of area
$\Delta C$	Increment of relative circulation
$\Delta \mathbf{r}$	Increment of the position vector of a point on a curve
$\Delta y$	Northward component of the increment of distance
$\frac{\partial v}{\partial n}$	Wind shear along the normal to the flow
$\zeta$	Scalar relative vorticity
$\zeta_a$	Scalar absolute vorticity
$\nabla_H$	Two-dimensional grad
$K_s$	Curvature of a streamline
$\lambda$	Coriolis parameter
$\mathbf{r}$	Position vector of a point on a curve
$\mathbf{v}$	Wind velocity
$v$	Wind speed
$v_y$	Scalar northward component of wind speed
$\phi$	Latitude
$\Omega$	Angular velocity of the earth



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## I. INTRODUCTION

The source of inspiration for this investigation was the desire to test, on a statistical basis, the frequently-used forecasting rule that wind flow from the south causes convergence and resulting convective weather phenomena. The subject has been of interest to other investigators. Namias and Clapp, on 10,000-foot maps, computed under steady state conditions "inertia trajectories"; i.e., trajectories of parcels of air moving such that the vertical component of the absolute vorticity remained constant [ 5 ]. The inertia trajectory was determined by the initial latitude of the parcel, as well as curvature, lateral shear, speed, and direction of the parcel. These vorticity trajectories were found to be not parallel to the isobars, presumably due to the modifying effects of convergence or divergence. If the isobars curved more cyclonically than the trajectories, convergence was present; if more anti-cyclonically, divergence existed. The fields of divergence and convergence were then related to areas of heating and cooling. The results of this research were applied to the trajectory method of forecasting by making qualitative corrections to the computed vorticity trajectories on the five-day mean 10,000-foot charts and thereby obtaining a more accurate prognosis of the future positions of troughs and ridges at that level.

Another paper related to the subject appeared in 1946, in which Houghton and Austin constructed charts of horizontal divergence [ 4 ]. One possible approach to the problem of computing quantitatively the time rate of change of vorticity would involve the construction and use of such charts, but as will be seen later, a different technique has been employed.



A paper somewhat more closely allied to the subject under consideration was published in 1950 [ 1 ]. Relative vorticity was computed at the 700-millibar level by the use of the expression for vorticity in natural coordinates, and correlated with the heights of the tops of cumuliform clouds over the ocean along the Washington-Bermuda airways route. The authors had considerable cloud data. Pilots flying the route were required to construct cross-sections showing weather encountered along the route upon completion of each flight. A large initial sample of 454 observations was obtained, as well as 210 additional observations later. In the original sample, the linear coefficient of correlation between vorticity and cloud-top height was 0.84; in the later sample, 0.65.

In a paper that appeared in March, 1951, H. T. Harrison applied qualitative synoptic tests to a group of forecasting rules [ 2 ].

Pertinent to this discussion is the rule:

A deep current of air moving northward across latitude lines in a straight or cyclonically curved path is undergoing convergence. Clouds and rainfall will develop which will be abundant if the airstream is curving cyclonically.

Harrison applied this rule to 39 actual cases, at the 700-millibar level, and found that 92% of them verified the rule.

The data in the above-mentioned papers was obtained from upper-level charts; in this paper, as will be explained later, mean sea level charts were used.





## II. THE THEORY OF RELATIVE VORTICITY

By way of review recall that the definition of vorticity came about in connection with the subject of the relative circulation about closed individual curves in a fluid. [ 3 ]. Relative vorticity (denoted by  $\zeta$ ) was defined as  $\lim_{\Delta A \rightarrow 0} \frac{\Delta C}{\Delta A}$ , and on the basis of this definition, the relative circulation about a closed curve  $C'$  was shown to be

$$C = \oint_{C'} \mathbf{v} \cdot d\mathbf{r} = \int_S \zeta \, dA$$

In rectangular coordinates, the vector representing relative vorticity is equal to the curl of relative velocity. In computing relative vorticity from actual weather maps, however, it was found more convenient to use the natural coordinate expression of this quantity;

$$\zeta = v K_s - \frac{\partial v}{\partial n} \quad (1)$$

where;

$v$  = wind speed

$K_s$  = curvature of the streamlines (positive for cyclonic curvature; negative for anti-cyclonic curvature)

$\frac{\partial v}{\partial n}$  = wind shear along the normal to the flow. If the normal,  $n$ , is considered to the left of the flow, the quantity  $\frac{\partial v}{\partial n}$  is negative for cyclonic shear, positive for anti-cyclonic.

The other basic equation pertinent to this development is the relative vorticity theorem [ 3 ]. The basic assumptions for this theorem were;

- (1) Frictionless fluid,
- (2) Nearly horizontal motions, and
- (3) Horizontal solenoids negligible.



By way of further illustration, let us suppose that the crust is composed of two layers of different materials, the upper layer being of thickness  $h$  and the lower layer of thickness  $h'$ . Let  $\rho$  and  $\rho'$  be the densities of the two layers, and let  $\mu$  and  $\mu'$  be the moduli of rigidity. Then the total stress  $P$  at the base of the crust is given by

$$P = \rho g h + \rho' g h' + \frac{\mu}{h} \epsilon + \frac{\mu'}{h'} \epsilon'$$

where  $\epsilon$  and  $\epsilon'$  are the strains in the two layers. If we suppose that the crust is in equilibrium, then the total stress  $P$  must be equal to the weight of the overlying material. This gives us the following equation:

$$\rho g h + \rho' g h' + \frac{\mu}{h} \epsilon + \frac{\mu'}{h'} \epsilon' = \rho g h + \rho' g h'$$

which simplifies to

$$\frac{\mu}{h} \epsilon + \frac{\mu'}{h'} \epsilon' = 0$$

or, if we suppose that the strain is the same in both layers, then we have

$$\frac{\mu}{h} + \frac{\mu'}{h'} = 0$$

which is the condition for equilibrium. This shows that the crust is in equilibrium when the total stress is equal to the weight of the overlying material. The above results are valid for a crust of uniform thickness. If the crust is of variable thickness, then the results are more complicated.

- (1) The crust is composed of two layers of different materials.
- (2) The crust is in equilibrium.
- (3) The total stress is equal to the weight of the overlying material.

By the use of the relationship that the time rate of change of absolute vorticity per unit absolute vorticity is equal to the negative horizontal divergence of the velocity; i.e.,

$$\frac{1}{\zeta_a} \frac{d\zeta_a}{dt} = -\nabla_H \cdot \mathbf{V}$$

The equation of relative vorticity was derived:

$$\frac{d\zeta}{dt} = -\frac{2\Omega \cos \phi}{a} v_y - (\zeta + \lambda) \nabla_H \cdot \mathbf{V} \quad (2)$$

where:

$\Omega$  = angular velocity of the earth ( $7.292 \times 10^{-5}$  radians per second)

$\phi$  = latitude of the parcel of air

$v_y$  = scalar northward component of the velocity of the wind

$\lambda$  = coriolis parameter

$\nabla_H \cdot \mathbf{V}$  = horizontal divergence of the velocity

Consider equation (2) for the case of southerly winds. If relative vorticity remains constant,  $\frac{d\zeta}{dt} = 0$ , and

$$\frac{2\Omega \cos \phi}{a} v_y = -(\zeta + \lambda) \nabla_H \cdot \mathbf{V} \quad (3)$$

For southerly flow,  $v_y$  is positive since the positive direction of  $y$  is taken as due north. Therefore, from equation (3), it is seen that horizontal convergence results. If vorticity is increasing,  $\frac{d\zeta}{dt} > 0$ , and from equation (2), horizontal convergence occurs. When vorticity is decreasing,  $\frac{d\zeta}{dt} < 0$ , which from equation (2) will occur either: (a) when the term  $(\zeta + \lambda) \nabla_H \cdot \mathbf{V} = 0$  (zero divergence); (b) when this same term is greater than zero (divergence); or (c), where this term is negative (convergence) but of smaller absolute value than  $\frac{2\Omega \cos \phi}{a} v_y$ .



Neglecting wind shear, these observations are easily depicted by a schematic diagram:

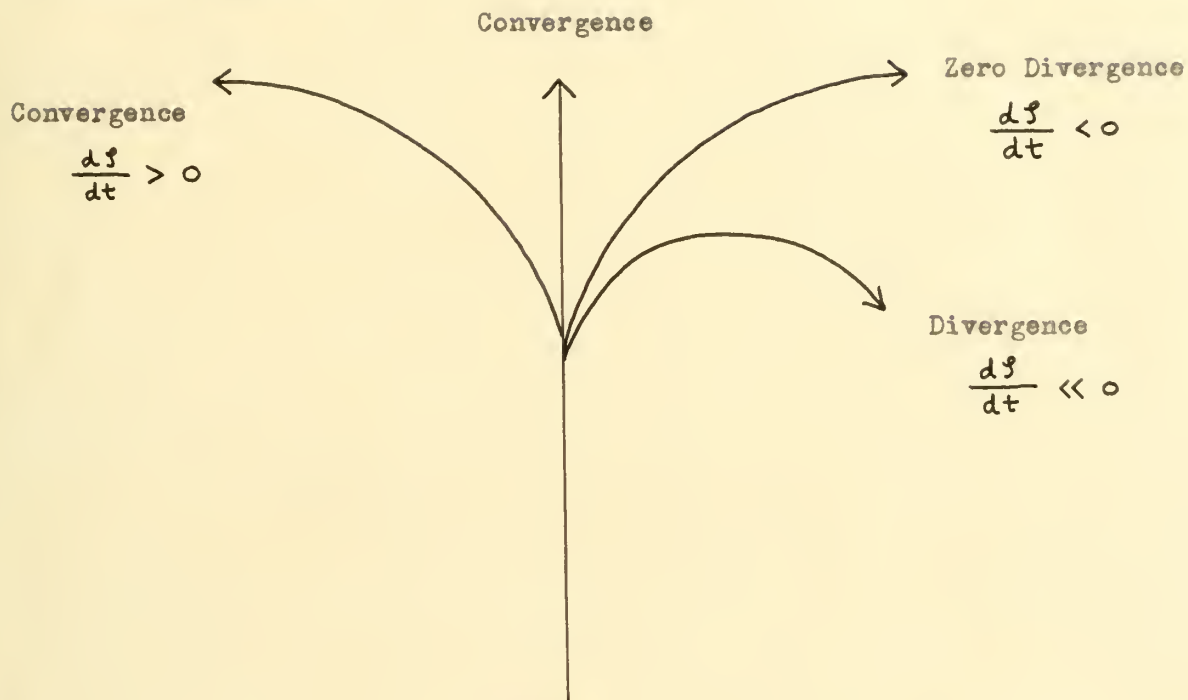


Figure 1.

#### Schematic Diagram Involving Southerly Winds

Now consider northerly flow, in which case  $v_y$  is negative. In straight northerly flow, it is apparent from equation (3) that divergence occurs. If vorticity is decreasing,  $\frac{d\mathcal{S}}{dt} < 0$ , and from equation (2), horizontal divergence results.  $\frac{d\mathcal{S}}{dt} > 0$  either when: (a), the term  $(\mathcal{S} + \lambda) \nabla_H \cdot \mathbf{V} = 0$  (zero divergence); (b), this term is less than zero (convergence); or (c) this term is positive (divergence) but less in absolute value than  $-\frac{2 \Omega \cos \phi v_y}{a}$ .





These remarks can be schematically represented, neglecting wind shear, as:

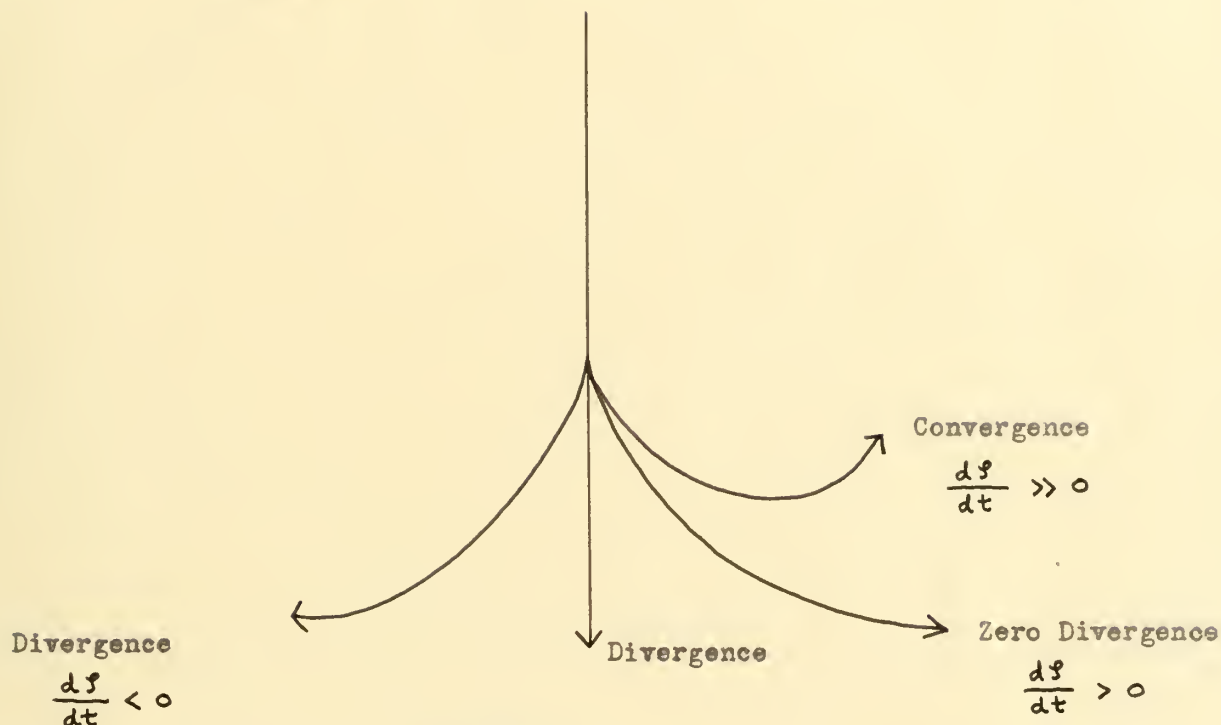
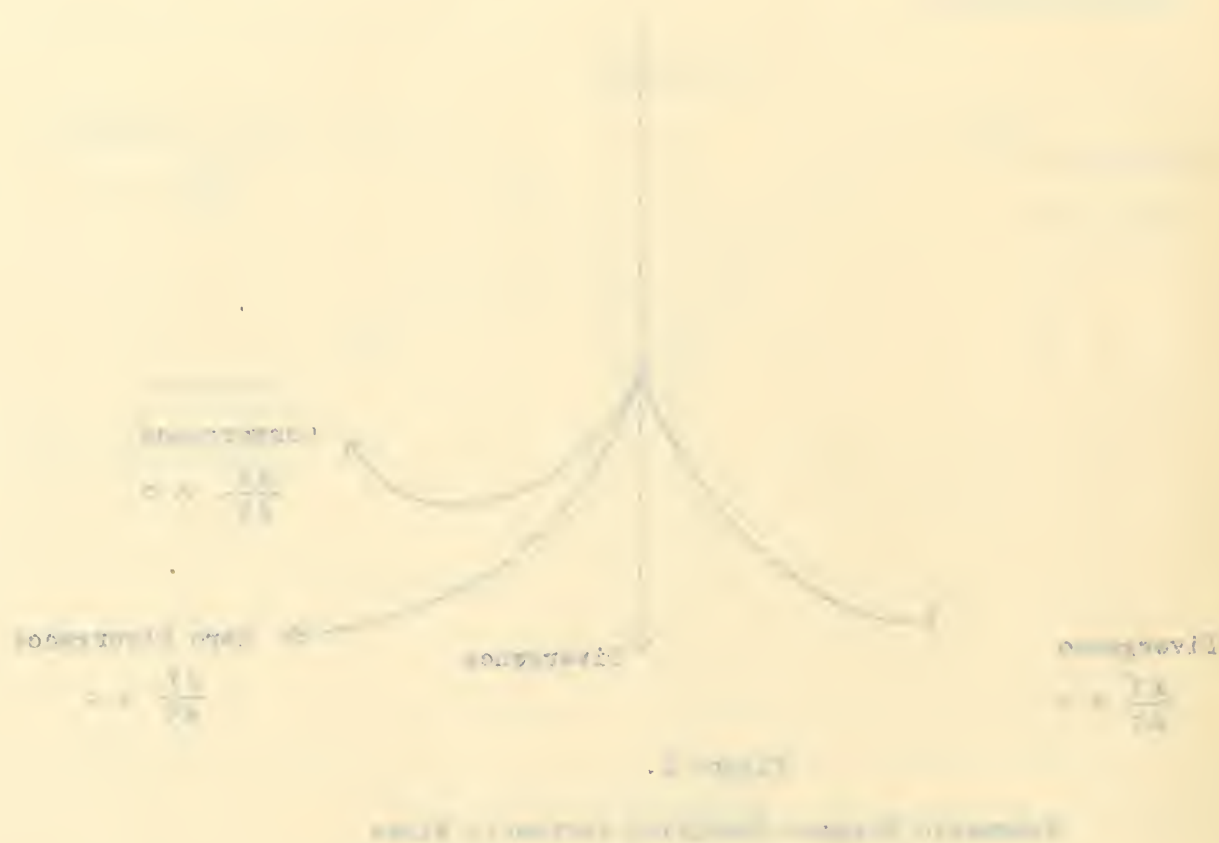


Figure 2.

#### Schematic Diagram Involving Northerly Winds

It has been tacitly assumed in the development above that the factor,  $\mathfrak{F} + \lambda$ , is positive. That this is generally true can be easily shown. The coriolis parameter,  $\lambda$ , is always positive in the northern hemisphere. At latitude  $30^\circ$  it has the smallest magnitude in the geographical region considered in this paper,  $0.26251 \text{ hours}^{-1}$ . In 268 computations of  $\mathfrak{F}$ , only in two instances of negative vorticity did the absolute value exceed 0.26251.





It has been shown that the curves  $y = x^2$ ,  $y = x^3$ , and  $y = x^4$  all pass through the origin and the point (1,1). It is also true that the curves  $y = x^5$ ,  $y = x^6$ , and  $y = x^7$  all pass through the origin and the point (1,1). In fact, for any positive integer  $n$ , the curve  $y = x^n$  passes through the origin and the point (1,1). This is because  $0^n = 0$  and  $1^n = 1$  for any positive integer  $n$ .

The preceding observations may be summarized as follows:

(1) Convergence occurs when there is straight southerly flow; divergence, when straight northerly flow.

(2) When  $\frac{d\zeta}{dt} > 0$ ; i.e., the flow is becoming less anti-cyclonic or more cyclonic, there is always convergence when the flow is southerly. There may be convergence or divergence when the flow is northerly, convergence occurring in cases where  $\frac{d\zeta}{dt}$  is of a relatively larger order of magnitude.

(3) When  $\frac{d\zeta}{dt} < 0$ ; i.e., the flow is becoming less cyclonic or more anti-cyclonic, there is always divergence when the flow is northerly; there may be divergence or convergence when the flow is southerly, divergence occurring when the absolute value of  $\frac{d\zeta}{dt}$  is of a relatively larger order of magnitude.

Hence, it is apparent from theoretical considerations that the blanket use of the forecasting rules of prognosticating convergence and convective activity with southerly flow and divergence with northerly flow is invalid. However, one would be led to expect a fairly high positive correlation between the rate of change of vorticity and convective activity, providing of course, that the model on which the theorem is based is sufficiently representative of average atmospheric conditions.

The following is a list of the names of the persons who have been elected to the office of the President of the United States since 1789.

(1) George Washington (1789-1797) (2) John Adams (1797-1801)

(3) Thomas Jefferson (1801-1809) (4) James Madison (1809-1817)

(5) James Monroe (1817-1825) (6) John Quincy Adams (1825-1829)

(7) Andrew Jackson (1829-1837) (8) Martin Van Buren (1837-1841)

(9) William Henry Harrison (1841-1845) (10) John Tyler (1845-1849)

(11) Zachary Taylor (1849-1850) (12) Millard Fillmore (1850-1853)

(13)

(14) Franklin Pierce (1853-1857) (15) James Buchanan (1857-1861)

(16) Abraham Lincoln (1861-1865) (17) Andrew Johnson (1865-1869)

(18) Ulysses S. Grant (1869-1877) (19) Rutherford B. Hayes (1877-1881)

(20) James A. Garfield (1881-1881) (21) Chester A. Arthur (1881-1886)

(22)

(23) Grover Cleveland (1886-1889) (24) Benjamin Harrison (1889-1893)

(25) William McKinley (1897-1901) (26) Theodore Roosevelt (1901-1909)

(27) William Howard Taft (1909-1913) (28) Woodrow Wilson (1913-1921)

(29) Warren G. Harding (1921-1923) (30) Calvin Coolidge (1923-1933)

(31) Herbert Hoover (1929-1933) (32) Franklin D. Roosevelt (1933-1945)

(33) Harry S. Truman (1945-1953) (34) Dwight D. Eisenhower (1953-1961)

(35) John F. Kennedy (1961-1963) (36) Lyndon B. Johnson (1963-1969)

### III. THE CORRELATION OF RELATIVE VORTICITY NEGLECTING WIND SHEAR WITH CONVECTIVE WEATHER PHENOMENA

The author has collected observations from surface weather maps in order to obtain a measure of the extent to which the theory applies in the atmosphere. The first phase of the investigation involved only a correlation of relative vorticity itself with convective weather phenomena. One hundred and twenty eight calculations were made of relative vorticity in cases where wind shear was negligible, with the use of equation (1). To simplify the calculations, only situations approximating steady-state conditions were used. The isobars which were considered to be streamlines were then also regarded as approximate trajectories of the air parcels. The curvature of the streamlines,  $K_s$ , was determined by comparison with a series of arcs drawn on opaque paper. The units of curvature used were nautical miles<sup>-1</sup>, and the velocity,  $v$ , estimated by use of a geostrophic wind scale, was measured in knots; i.e., nautical miles hours<sup>-1</sup>. Thus the units of relative vorticity were hours<sup>-1</sup>.

The average relative vorticity was computed for a square bounded by five degrees of latitude and five degrees of longitude. Such squares were considered only in the Great Plains region of the United States, generally from 30 to 55 degrees north latitude, and from 90 to 100 degrees west longitude. It was deemed inadvisable to consider areas where convection might be the result of orographic effects. Attention was also given to avoiding such cases where convection was the result of frontal lifting.

A "convective index" for the square area was then also computed. This was accomplished by assigning various weights to the different types of



The subject was collected observations in a series of cases.

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clouds and other weather phenomena, higher weights being assigned to the weather types indicative of greater convective activity. Standard numerical symbols for clouds and weather types were used, as designated by the U. S. Weather Bureau [7]. The assigned weights are shown below:

Table 1.

Weight	N	C <sub>L</sub>	C <sub>M</sub>	C <sub>H</sub>	WW
10		9			17 29 65 75 82 90-99
8		2,3,7	8		50-59 60-64, 66-69 70-74, 76-79 80-81, 83-89
6					14-16 20-27
4		1,4,5,8	6,9		
2			3,4,5,7	9	
0	0	6	1,2	1-8	11 12 28 40-49

Table of Weights Assigned to Weather Phenomena





The convective index was computed by selecting from each reporting station in the square the symbol with the greatest weight and determining the average of these weights. The linear correlation coefficient was then obtained between the convective indices and the average relative vorticities, rounded off to the nearest hundreth and multiplied by  $10^2$  for the purpose of convenient units. Separate correlations were made for the cases of southerly winds, of northerly winds, and then of all cases combined. The scatter diagrams for these correlations and the results are shown on the succeeding pages.

The committee found that the evidence presented by the witnesses was not sufficient to establish the facts of the case. The committee also found that the witnesses were not credible and that the evidence was not reliable. The committee therefore recommended that the charges against the accused be dropped.

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TABLE 2. SCATTER DIAGRAM OF RELATIVE VORTICITY AND CONVECTIVE INDEX  
(SOUTHERLY WINDS)

		Convective Index										Totals
$v_j$	$u_i$	-4	-3	-2	-1	0	1	2	3	4	5	
-8												0
-7				2								2
-6												0
-5				1	1	1						3
-4		2		3								5
-3		2	3	5	3	1	3	1		2		20
-2		5	4	3	6	6	3	3		1	1	32
-1		5	1	3	1	2	1	1	1			15
0		3		4	2	2		1				12
1					1	1						2
2						1						1
3												0
Totals		17	8	21	14	14	7	6	1	3	1	92



TABLE 3. SCATTER DIAGRAM OF RELATIVE VORTICITY AND CONVECTIVE INDEX  
(NORTHERLY WINDS)

		Convective Index										Totals
$v_j$	$u_1$	-4	-3	-2	-1	0	1	2	3	4	5	
-16						1						1
-15												0
-14												0
-13												0
-12												0
-11												0
-10												0
-9												0
-8												0
-7												0
-6												0
-5			2					1				3
-4		2				1	2					5
-3		3	1									4
-2		3	2	1	1							7
-1		2	1									3
0		1	2				1		1	2		7
1								1				1
2			1							1		2
3								1				1
4										1		1
5												0
6												0
7												0
8												0
9										1		1
Totals		11	9	1	1	2	3	3	1	5	0	36

$y \times 10^2$  Hours<sup>-1</sup>





TABLE 4. SCATTER DIAGRAM OF RELATIVE VORTICITY AND CONVECTIVE INDEX  
(COMBINED SOUTHERLY AND NORTHERLY WINDS)

		Convective Index										Totals
$v_1$	$u_1$	-4	-3	-2	-1	0	1	2	3	4	5	
-16						1						1
-15												0
-14												0
-13												0
-12												0
-11												0
-10												0
-9												0
-8												0
-7				2								2
-6												0
-5			2	1	1	1		1				6
-4		4		3		1	2					10
-3		6	4	5	3	1	3	1		2		24
-2		8	6	4	7	6	3	3		1	1	39
-1		7	2	3	1	2	1	1	1			18
0		4	2	4	2	2	1	1	1	2		19
1					1	1		1				3
2			1			1				1		3
3								1				1
4										1		1
5												0
6												0
7												0
8												0
9										1		1
Totals		28	17	22	15	16	10	9	2	8	1	128

3 x 10<sup>2</sup> Hours



Table 5.

Wind Flow	Average Vorticity	Range of Vorticity	Average Con- vective Index	Range of Con- vective Index	r
Southerly Winds	-.02005	-.12987 to .05560	3.310	.00 to 9.33	.0232
Northerly Winds	-.01283	-.30000 to .20850	3.328	.00 to 8.00	.3408
Combined	-.01802	-.30000 to .20850	3.315	.00 to 9.33	.2744

Results of Correlation of Vorticity and Convective Index

Table 1

Year	Number of migrants	Number of returnees	Number of refugees	Number of asylum seekers	Number of naturalized citizens
1990	10,000	5,000	15,000	20,000	10,000
1991	12,000	6,000	18,000	22,000	12,000
1992	15,000	8,000	20,000	25,000	15,000

Source: Ministry of Foreign Affairs and International Trade

Another paper has been prepared in conjunction with this study, correlating convection with relative vorticity where wind shear was included [6]. In his research, Lieutenant Tatone made 140 computations of convective indices and relative vorticity. The same measuring techniques were employed, and a linear correlation coefficient of  $-0.1050$  for all cases combined was obtained.

From these computations it is seen that vorticity itself did not correlate well with convective activity. In the Washington-Bermuda airways study, better correlations were obtained;  $0.84$  in the sample of 454 observations and  $0.65$  in the sample of 210 [1]. A possible explanation for these higher correlations, aside from the fact that the samples were larger, is that vorticity at the 700-millibar level is perhaps more representative for the layer above the earth's surface than vorticity at the surface itself. Also, the height of cloud tops might be a better indication of convective activity than surface weather observations.





#### IV. THE CORRELATION OF THE TIME RATE OF CHANGE OF RELATIVE VORTICITY WITH CONVECTIVE WEATHER PHENOMENA

The final phase of this investigation involved the correlation of the time rate of change of relative vorticity with convective weather phenomena, under nearly steady-state conditions. By dividing equation (2) by  $v_y$  is obtained the relationship:

$$\frac{d\zeta}{v_y dt} = \frac{d\zeta}{dy} = - \frac{2\omega \cos \phi}{a} - \frac{\zeta + \lambda}{v_y} \nabla_h \cdot \mathbf{v}$$

The quantity,  $\frac{d\zeta}{dy}$ , was computed by the use of finite differences:

$\frac{\Delta \zeta}{\Delta y} = \frac{\zeta_2 - \zeta_1}{\Delta y}$ . Two adjacent square areas were selected such that the air parcel could be considered as moving along the streamlines from the center of the first square to the center of the second. At the center position of each square, the parcel was regarded as possessing the average relative vorticity for that square. Since the squares were adjacent and five degrees of latitude in size, the movement of the parcel,  $\Delta y$ , was 300 nautical miles, positive for southerly winds, negative for northerly winds.

Vorticity was determined again by the use of equation (1); however, in this phase of the work wind shear was included. In measuring the shear, the direction of the normal was regarded to the left of the direction of flow. The term  $\frac{\partial v}{\partial n}$ , or  $\frac{\Delta v}{\Delta n}$ , was then negative when the flow decreased along the normal (cyclonic shear), positive when it increased (anti-cyclonic shear). Then the two terms of equation (1) were added algebraically to obtain  $\zeta$ , and the rate of change of  $\zeta$  determined from the relationship:  $\frac{\Delta \zeta}{\Delta y} = \frac{\zeta_2 - \zeta_1}{\Delta y}$ , in hours<sup>-1</sup> nautical miles<sup>-1</sup>.

# IV. THE COMPARISON OF THE TWO METHODS OF MEASURING THE EFFECT OF THE TEMPERATURE ON THE RATE OF REACTION

The first method of measuring the rate of reaction is the method of initial rates.

The second method of measuring the rate of reaction is the method of integrated rates.

Comparison of the two methods of measuring the rate of reaction is made as follows:

Let  $k_1$  be the rate constant for the reaction:

$$A + B \rightarrow C + D$$

The rate of reaction is given by the equation:

$$-\frac{d[A]}{dt} = -\frac{d[B]}{dt} = \frac{d[C]}{dt} = \frac{d[D]}{dt}$$

Let  $k_2$  be the rate constant for the reaction:

$$A + B \rightarrow E + F$$

The rate of reaction is given by the equation:

$$-\frac{d[A]}{dt} = -\frac{d[B]}{dt} = \frac{d[E]}{dt} = \frac{d[F]}{dt}$$

Let  $k_3$  be the rate constant for the reaction:

$$A + B \rightarrow G + H$$

The rate of reaction is given by the equation:

$$-\frac{d[A]}{dt} = -\frac{d[B]}{dt} = \frac{d[G]}{dt} = \frac{d[H]}{dt}$$

Let  $k_4$  be the rate constant for the reaction:

$$A + B \rightarrow I + J$$

The rate of reaction is given by the equation:

$$-\frac{d[A]}{dt} = -\frac{d[B]}{dt} = \frac{d[I]}{dt} = \frac{d[J]}{dt}$$

Let  $k_5$  be the rate constant for the reaction:

$$A + B \rightarrow K + L$$

Let  $k_6$  be the rate constant for the reaction:

The convective index for each square was computed according to the method described previously; i.e., of calculating the average of the highest-weighted phenomena reported by the stations in that area. Then the indices of the two squares were used to arrive at an average convective index associated with the movement of the parcel. The same geographical area was used and again instances where frontal activity was present were avoided.

For convenience in units, the scatter diagram was based on  $\frac{d\mathcal{I}}{dy} \times 10^4$ . It is shown below. Since only 9 of the 70 observations involved northerly winds, all cases were combined on a single scatter diagram.

A coefficient of 0.0513 was obtained, representing the linear correlation between the convective index and the rate of change of relative vorticity. The average convective index was 2.952 and the average rate was 0.0000424 hours<sup>-1</sup> nautical miles<sup>-1</sup>.

The method for determining the relative amounts of the  
 various components in a mixture is by measuring the area under  
 each peak in the chromatogram. This is done by integrating  
 the peaks or by using a planimeter. The area under each peak  
 is proportional to the amount of that component in the mixture.  
 The relative amounts of the components are then calculated by  
 dividing the area under each peak by the total area under all  
 the peaks.

The concentration of a component in a mixture can be determined  
 by comparing the area under its peak in the chromatogram to  
 the area under a peak of a known concentration of the same  
 component. This is done by using a standard solution of the  
 component. The concentration of the component in the mixture  
 is then calculated by dividing the area under its peak by the  
 area under the peak of the standard solution and multiplying  
 by the concentration of the standard solution.



TABLE 6. SCATTER DIAGRAM OF RATE OF CHANGE OF RELATIVE VORTICITY AND CONVECTIVE INDEX

[illegible]









## V. RESULTS AND INTERPRETATIONS

On the basis of the observations and measurements made in this investigation it appears that under conditions approximating those of steady-state neither vorticity nor the rate of change of relative vorticity correlates well with convective activity. One pertinent factor, not considered, is the moisture content of the air. However, since the wind flow in the majority of the cases was from the south, and had a long trajectory over the Gulf of Mexico, the moisture present was assumed to be adequate for the production of weather, given sufficient convection. It may also be true that the layer of air as represented by sea level isobars, extending from the surface to about 5,000 feet, is not sufficiently deep for testing the theory. Perhaps the layer from the surface to 500 or 700 millibars would give a better correlation. It is also possible that the assumptions on which the relative vorticity theorem is based are too restricted and do not represent average atmospheric conditions. Finally, the measuring techniques may have been too crude; however, the quantities that were involved in this work were measured as accurately as possible from maps available in the normal aerological office.

## 1. CONCLUSIONS AND INTERPRETATION

On the basis of the observations and measurements made in this investigation it appears that under conditions approximating those of steady-state motion velocity is the type of change in relative velocity concerned with the acceleration activity. The present theory, not consistent, is the relative content of the air. However, there was time in the majority of the cases was from the source, and had a look at the very end of the air, the relative pressure was reduced in the direction for the direction of motion, given suitable conditions. It may also be seen from the figure of air as represented by the figure, according to the figure is about 2,000 feet, is not sufficient for the motion of the air. Perhaps the report from the source to the 1000 miles would give a better impression. It is also possible that the assumption of which the relative velocity between is made was not restricted and the air pressure between atmospheric conditions. Finally, the measuring technique was not very good; however, the conditions that were involved in this work were regarded as satisfactory as possible from more available in the normal vertical direction.

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